# Sub watershed Prioritization Based on Morphometric Techniques and Machine Learning Approaches in The Sobha River Basin Using Geospatial Techniques, India

<sup>1</sup>Mousumi Roy, Prof. <sup>2</sup>Sujit Mandal

<sup>1</sup>Research Scholar, <sup>2</sup>Professor

<sup>1,2</sup> Department of Geography, Diamond Harbour Women's University, South 24 Parganas, India<sup>1</sup>

#### **Abstract:**

Morphometric analysis is important in any hydrological investigation and it is inevitable in the development and management of drainage basin. In the present paper, an attempt has been made to study the morphometric parameters of the Sobha micro-watershed in Purulia, West Bengal. To know the watershed dynamics and manage those in terms of drought and flash flood reduction of a plateau river (Sobha River) and to manage soil erodibility and groundwater potentiality; we have taken the aforesaid watershed for our case study. For this study, hydrology-based GIS functions have been used in the evaluation of linear, aerial, and relief morphometry of the Sobha watershed. Several 11 morphometric indices have been generated for all the subwatersheds to understand the geomorphological behavior of the Sobha watershed. The total area of the basin is 137 Sq.km and lies between 22°42'19"N to 23°42'00"N latitude and 85°49'19"E to 86°54'25"E longitude. Poor soil cover, sparse vegetation, low amount of rainfall, and lack of soil moisture characterize the study area for most of the year. So the entire study area has been further divided into 37 sub-watersheds, named MWS 1, MWS 2.....MWS 37 ranges in several geographical areas. The drainage density values of sub-watersheds MWS 10, 14, 16, 23, and 24 indicate that it has highly resistant power with dense vegetative cover and high relief. The elongation ratio varies from 12.44 to 77.89 which indicates very rugged terrain and steep ground slope. The composite score values are calculated and the sub-watersheds 7, 15, 16, 21, 23, 27, 34 and 36 has a minimum composite score value of 1, 2, 5, 11, 12, 20, 27, 28, 30, and 31 is likely to be subjected to maximum soil erosion and susceptible to natural hazards. Hence this maximum vulnerable watershed should be provided with immediate soil conservation techniques.

**Index Terms:** Morphometric Analysis, Geospatial techniques, Machine learning, Sub-watersheds Prioritization

#### 1. Introduction:

In fluvial landscape dynamisms are a common element and its determinants are geomorphology, geology, hydrology, soil, vegetation, climate, etc. of that region (Strahler 1964; Mesa 2006; Rekha et al. 2011; Romohoo et al. 2012; Puno and Puno 2019; Singh et al. 2021). Physical and hydrological characteristics like area, slope, size, shape, length, drainage density, etc. are correlated to understand the control of drainage basin morphometry upon the landscape of a basin (Clarke 1966; Rastogi and Sharma 1976; Rastogi et al. 1976; Pakhmode et al. 2003; Sreedevi et al. 2009; Das et al. 2013; Amulya et al. 2018). This analysis is very useful for any kind of basin-scale management activities such as soil erosion, flood modeling, sustainable resource management, basin evolution prediction, etc. So, this method is a useful alternative tool where field data are limited because of accessibility, remoteness or scarcity of field based technological advancement (Nookaratnam et al. 2005; Borga et al. 2008; Javed et al. 2009; Sharma et al. 2010; Romshoo et al. 2012; Altaf et al. 2013; Aher et al. 2014; Ajay et al. 2014; Bhatt and Ahmed 2014; Abuzied et al. 2016; Ayele et al. 2017; Balasubramanian et al. 2017; Gajbhiye and Sharma 2017; Taha et al. 2017; Ameri et al. 2018; Kannan et al. 2018; Shivhare et al. 2018; Andan et al. 2019; Asfaw and Workinch 2019; Charizopoulos et al. 2019; Gunjan et al. 2019; Hussein et al. 2019; Karabulut and Ozdemir 2019; Mahmood and Rahman 2019; Nitheshnirmal et al. 2019; Puno and Puno 2019; Abdeta et al. 2020; Alam et al. 2020; Arefin et al. 2020; Das 2020; Gabriel et

al. 2020; Meshram et al. 2020; Ogarekpe et al. 2020; Pan et al. 2020; Rajsekhar et al. 2020; Sangma and Guru 2020; Obeidat et al. 2021; Singh et al. 2021, Tukura et al. 2021). This technique is also very helpful for developing and underdeveloped countries where food security strategy is of utmost important and thus water resource conservation is needed to protect the basin environment (Desa et al. 2005; Sharma et al. 2005; German et al. 2007; Kerr 2007; Draghouth et al. 2008; Kumar and Palanisami 2009; Alemu and Kidane 2014; Igbal and Sajjad 2014; Mekhonnen and Fekadu 2015; Worku and Tripathi 2015; Abdeta et al. 2020).

The morphometric analysis involves numerical measurement of size, shape, and other related dimensions of the basin surface (Clarke 1996; Agarwal 1998; Obi et al. 2002). It involves numerous quantitative parameters which are practices nowadays are congregated from different pioneer works like Horton, Smith, Strahler; Miller and Schumm (Horton 1932, 1945; Smith 1950; Strahler 1952, 1964; Miller 1953; Schumm 1956; Obi Reddy et al. 2002; Vaidya et al. 2013; Kaur et al. 2014). This approach measured various attributes in three broad categories such as linear, areal, and relief aspects to get insights about the basin characteristics (Melton 1957; Strahler 1964; Sreedevi et al. 2009; Arnous et al.2011; Charizopoulos et al.2019; Abedeta et al. 2020; Obeidat et al. 2021; Singh et al. 2021). Morphometry-based prioritization of different sub-watersheds is made by comparing each of them (Sharma et al. 2010; Kumar et al. 2015; Kumar and Lal 2017; Thapliyal et al. 2017).

The Sobha River basin is characterized as semi-arid to sub-humid in different seasons throughout all the year. As most of the inhabitants of this plateau region depend on limited agriculture and livestock farming they depend on groundwater resources. Geologically the basin region is typically a rolling plain, barren land, and granite-gneissic-quartzite complex zone. Thus, scarcity of surface water is most common. On the other hand, dependency on groundwater resulted in the declination of groundwater levels for unrestricted and overutilization of this resource. So, reaching a scale watershed prioritization scheme is a requisite of the region. For this reason, we used SRTM DEM data to investigate different morphometric aspects (Moore et al. 1991; Nag 1998; Farr and Kobrick 2000; Smith and Sandwell 2003; Grohmann 2004; Chopra et al. 2005; Das and Mukherjee 2005; Korup et al. 2005; Ratnam et al. 2005; Grohmann et al. 2007; kale and Shejwalker 2007; Lindsay and Evans 2008; Rudraiah et al. 2008; Sreedevi et al. 2009; Patel and Sarkar 2010; Wang et al. 2010; Pareta and Pareta 2011; Altaf et al. 2013; Magesh et al. 2013; Jacques et al. 2014; Das et al. 2016; Senthamizhan et al. 2016). Remote sensing techniques have great potential because of wide synoptic view, multispectral and multi-temporal capabilities, repetitiveness, and compute-compatibility (Pareta and Pareta 2011, 2012; Das et al. 2018). Hence integrated management of river basins remote sensing and GIS-based geospatial techniques are reliably used in present-day studies (Fransisco and Rola 2004; Pandey et al. 2007; Chatterjee et al. 2013; Chandniha and Kansal 2014; Okumura and Araujo 2014; Syed et al. 2017; Rahmati et al. 2019; Abedeta et al. 2020; Borah and deka 2020; Obeidat et al. 2021). Remote sensing and GIS tools are utilized for the water resource management and development of water resources. Due to advancements in satellite and sensing technology, it is now possible to map finer details of the earth's surface and it provides scope for micro-level planning and management. The present study aims at the proper management of water resources and controlling surface soil loss. Water resource management by prioritization of micro-watershedbased analysis using remote sensing data and GIS overlying techniques. This research contributes to the framework for management practices of basin as micro-watersheds wise implemented by local administrative authorities as well as homogeneous basin management programs. This study is also helpful for increasing the agriculturally based livelihood, and irrigation facilities and to find the solution of uncontrolled soil loss. A watershed is an ideal unit for the management of water for land and water resources for mitigation of the impact of natural disasters for achieving sustainable development.

The significant factors for planning and development of a watershed are its physiography, drainage, geomorphology, soil, land use, land cover, and available water resources. Remote sensing and GIS are the most proven tools for watershed development as well as management and the studies on prioritization of micro-watershed development and management.

### 2. Study Area:

The study area is in Purulia, the westernmost district of the Indian state of West Bengal and a part of Jharkhand state in Chotanagpur Plateau is located between the graticule 22°42'19" N to 23°42'00"N latitude and 85°49'19"E to 86°54'25"E longitude covering an area of 166 sq km. The study area comprises with Sobha watershed of Purulia district of West Bengal, India. The Sobha River is a tributary of the Kistobazar River in the southern part of the Ajodhya hill which further moves toward south west and meets the Subarnarekha River.

It is traversed by the tropic of cancer. This is located in the foothill of Ajodhya in Baghmundi P.S. and the main areas are Ajodhya, Baghmundi and Mathaburu etc. The Bay of Bengal and the Hooghly estuary are within 200 km from the center of the district. It is surrounded by Paschim Medinipur, Bankura, and Burdwan districts of West Bengal and Dhanbad, Bokaro, Hazaribag, East and West Singhbhum of Jharkhand state. Its physiographic location is also distinguishable as a zone of transition between the young alluvial plains of West Bengal and the ancient plateau of Southeast Bihar.

Climatically, the area is sub-tropical and sub-humid, with hot wet summers and cool dry winters characterized by high evaporation and low precipitation with an annual mean temperature of 25.6°C and mean summer and mean winter temperature of 29.0°C and 21.3°C, respectively. The monsoon is the main source of precipitation, which shapes a 'funnel' like meteorological functions. It funnels the tropical monsoon current from the Bay of Bengal to the sub-tropical parts of North West India. It starts in May and continues up to October. It has an annual average precipitation of 1393 mm. About 82% of the annual rainfall occurs during the monsoon which lasts roughly from June to September but uneven, scanty, and erratic rainfall results in agricultural drought in the kharif season.



Plate 1: Some glimpse of the study area and different field observations

The main rivers passing through the district are Kangsabati, Kumari, Darakeswar, Subarnarekha, and Damodar but the study area comprises Turga Nala, Kistobazar Nala, and Daurigarha Nala. Soil erosion is the most prominent phenomenon in this area resulting huge deposition of fertile soil in the valley region. As a result, the area is facing a crisis due to the depletion of top fertile soil and water loss. The study area is a 'White Zone' concerning groundwater status (i.e., 60% of available Ground annual recharge is in use). In general, during the rainy season, the water table in the wells rises from 1.00 to 3.50 m bgl till the end of October and gradually falls to a maximum of 6 to 14 m bgl during April-May.

Geologically the area is a part of the Chotanagpur Gneissic complex of the Eastern Indian peninsular shield, lying to the north of Singhbhum craton. China clay occurrences in this area are invariably associated with granitic rocks and meta-sediments of the Chotanagpur Gneissic complex of the Precambrian age. The topography is undulating with moderate to gentle slopes. It has a thick stratigraphic succession of mostly Archaean granitic gneiss and to a much lesser extent, Quaternary semi-consolidated sediments, promocarboniferous sandstone, and shale. Precambrian massive granites and quartzite sand with recent alluvium sediments deposition. Mineralogically these rocks are composed mainly of quartz, feldspar, muscovite, biotite, illite, and kaolinite.

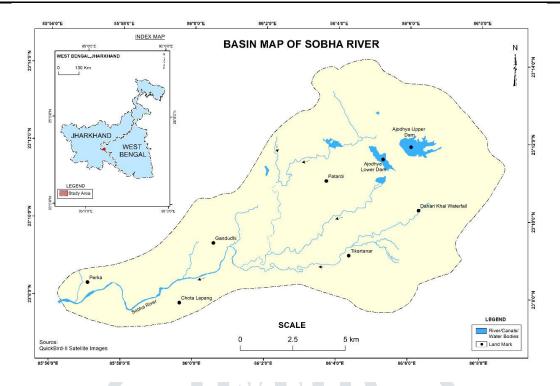


Figure 1: Location Map

# 3. Database and Methodology:

The present study is mainly concerned with evaluating morphometric characteristics of the river basin at the sub-watershed level to assess the prioritization of the Sobha micro-watershed with geospatial techniques. For this, both SOI (Survey of India) Toposheets (73I/4, 73E/16) and SRTM DEM data (30m. resolution) were used in the GIS environment. At the outset, SRTM data is used as input in Arc SWAT and delineates watersheds as well as 37 sub-watersheds. Georeferencing of toposheets was done in TNT Mips 2014 and ARC GIS 10.1 version software and mosaic the same to subset the study region.

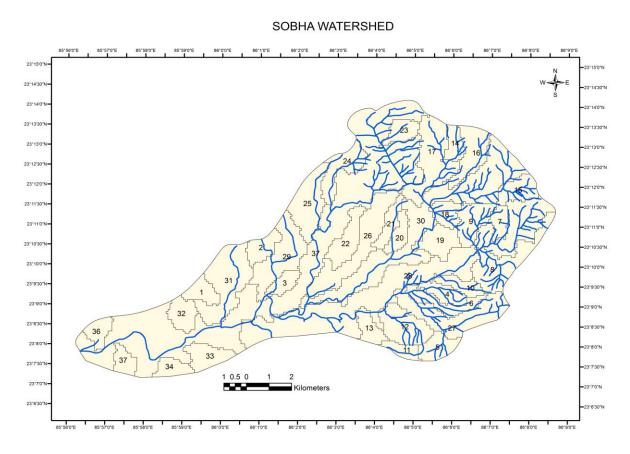


Figure 2: Sub-watersheds of the study area

The rectification method has been employed to obtain a geometrically corrected subset image. Standard methods of Strahler's, Horton's, Miller's, and Schumm's were applied to examine morphometric characteristics. Every single morphometric element is considered in the prioritization of watershed and weightage (composite score) has been assigned. For the linear aspect, high weightage was assigned for high values, and aerial aspect low weightage was assigned for high values (Panhalkar and pawar, 2011; Panhalkar et al., 2012) The compound values of all parameters at the sub-watershed level were calculated and highest priority has assigned to sub-watershed having lowest average compound weightage (composite score) and vice versa. Here priority implies the necessity of land resource conservation and the need to implement suitable remedial measures.

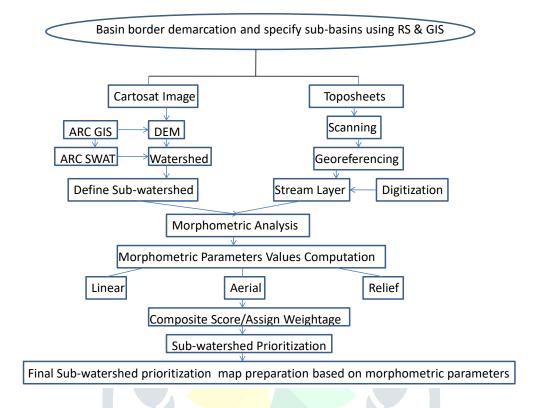


Figure 3: Methodology

# 4. Result and Discussion:

The prioritization of the Sobha watershed has been carried out based on morphometric characteristics. 37 sub-watersheds have been generated in ARC SWAT software. Details of morphometric analysis of the Sobha watershed are enumerated in Tables 1, 2, and 3.

**Table 1:** Descriptive Characteristics of morphometric parameters

Aspect	Parameters	Characteristics
A. Linear	Average length of stream (L <sub>um</sub> )	It is the characteristic property related to the drainage network and its associated surfaces. Generally, the higher the order, the longer the length of the stream is noticed in nature.
	Bifurcation Ratio (R <sub>b</sub> )	Its ratio characteristically ranges between 3.0 and 5.0 for watersheds in which the geologic structures do not distort the drainage pattern but if it is <1 then it is vice versa.
	Stream Length Ratio     (L <sub>ur</sub> )	It is a successive stream order which varies due to differences in slope and topographic conditions and has an important relationship with the surface flow discharge and erosional stage of the basin.
	• Stream Frequency (F <sub>s</sub> )	It is the total no. of stream segments of all orders per unit area. Generally, high stream frequency is related to impermeable sub-surface material, sparse vegetation, high relief conditions, and low infiltration capacity.
B. Aerial	• Length of Overland Flow (L <sub>g</sub> )	It is defined as the length of the runoff of the rainwater on the ground surface before it is localized into definite channels.
	Drainage Density (D <sub>d</sub> )	It is a measure of the degree of fluvial dissection. It indicates the closeness of spacing of channels influenced by resistance to erosion, infiltration capacity, vegetation cover, surface roughness and runoff intensity, climatic condition, etc. (Reddy et al., 2004) Low Dd leads to coarse drainage texture while high Dd leads to fine drainage texture.
	Constant of Channel maintenance (C)	The constant indicates the no. of km <sup>2</sup> of basin surface required to develop and sustain a channel 1 km long. The constant of channel maintenance indicates the relative size of landform units in a drainage basin and has a specific genetic connotation (Strahler, 1957).
	• Form Ratio (F <sub>f</sub> )	It is defined as the ratio of basin area to the square of basin length. The value of Ff would always be less than 0.7854 (for a perfectly circular basin). The smaller the value of Ff, the more elongated will be the basin. The basins with high Ff have high peak flows of shorter duration, whereas, elongated form factors have lower peak flows of longer duration.
	• Elongation Ratio (R <sub>e</sub> )	Elongation Ratio is defined as the ratio of the diameter of a circle of the same area as the basin to the maximum basin length. It helps to give an idea about the hydrological character of a drainage basin.
C. Relief	• Relief Ratio (R <sub>h</sub> )	The elevation difference between the highest and lowest points on the valley floor of a sub-watershed.
	• Ruggedness Number $(R_n)$	It is the product of the basin relief and drainage density and usefully combines slope steepness with its length.

 Table 2: Ranges of Morphometric Parameters

Aspects	<b>Morphometric Parameters</b>	Formula	Reference	Results (Ranges)
	Stream length Ratio (Lur)	$L_{ur}=L_{u-1}$	Strahler (1964)	0.691 - 7.154
<u>.</u>	Bifurcation ratio (R <sub>b</sub> )	$R_b = N_u / N_{u+1}$	Schumm (1956)	0.500 - 2.000
Linear	Total stream length (Lu)km	$L_u = L_{1+} L_{2+\dots +} L_n$	Strahler (1964)	1235.100 – 42747.470
1	Stream frequency (F <sub>s</sub> )	$F_s=N_u/A$	Horton (1932)	0.00000007 -
				0.00000186
	Drainage density (D <sub>d</sub> ) km/ km <sup>2</sup>	$D_d = L_u/A$	Horton (1945)	0.000162 - 0.202171
Aerial	Form ratio (F <sub>f</sub> )	$F_f=C_L/P$	Horton (1932)	121.4687 – 4765.1305
	Elongation ratio (R <sub>e</sub> )	$R_e=2/L_b*(A/\pi)^{0.5}$	Schumm (1956)	12.440 – 77.890
	Length of overland flow (Lg)	$L_g=A/2*L_u$	Horton (1945)	230.280 – 3089.000
	km			
	Constant of channel	$C=1/D_d$	Schumm (1956)	460.560 – 6177.990
	maintenance(C) km/ km <sup>2</sup>			
Relief	Relief ratio (R <sub>h</sub> )	$R_h=H/L_b$	Schumm (1956)	0.190 – 17.990
Re	Ruggedness number (R <sub>n</sub> )	$R_n = D_d * (H/1000)$	Strahler (1950)	0.024979 - 0.623150

 Table 3: Descriptive Analysis of Components

MWS 5	MWS 4	MWS 3	MWS 2	MWS 1	Micro-watershed
					code no.
	1823.67	1683.74	1404.88	1607.78	Average stream length
	1.000	1.000	1.000	1.000	Bifurcation ratio
	1.000	1.000	1.000	1.000	Stream length ratio
	0.000712	0.001607	0.001008	0.001237	Drainage density
	1403.52	622.09	992.11	808.25	Constant of channel maintenance
	701.76	311.04	496.06	404.12	Length of overland flow
0.00000068	0.00000039	0.00000095	0.00000072	0.00000077	stream frequency
	197.9775	2811.8941	4765.1305	2884.5362	Form ratio
	15.88	59.83	77.89	09.09	Elongation Ratio
	2.43	1.50	1.64	1.41	Relief ratio
	0.196648	0.046617	0.028223	0.037117	Ruggedness no.
510.29021	119.08568	294.05665	489.80438	309.44899	Composite Score

MWS 19	MWS 18	MWS 17	MWS 16	MWS 15	MWS 14	MWS 13	MWS 12	MWS 11	MWS 10	MWS 9	MWS 8	MWS 7	MWS 6
3205.34	1235.10	4495.30	4854.31	1522.35	1867.82	1934.40	2658.57	2658.57	1823.67	2876.22	2416.51	1496.78	3933.11
2.000	1.000	1.000	2.000	1.000	1.000	1.000	1.000	1.000	1.000	2.000	1.000	1.000	1.000
5.317	1.000	0.691	1.700	1.000	1.000	1.000	1.000	1.000	1.000	1.106	1.000	1.000	5.635
0.001519	0.000928	0.000933	0.001824	0.000899	0.001636	0.001574	0.000353	0.000353	0.000712	0.001291	0.000869	0.001408	0.002171
658.38	1077.37	77.1701	548.14	1111.92	611.11	635.20	2835.19	2835.19	1403.52	774.74	1150.38	710.18	460.56
329.19	538.69	535.89	274.07	555.96	305.56	317.60	1417.60	1417.60	701.76	387.37	575.19	355.09	230.28
0.00000142	0.00000075	0.00000042	0.00000113	0.00000059	0.0000008	0.00000081	0.00000013	0.00000013	0.00000039	0.00000135	0.00000036	0.00000094	0.00000110
940.2269	391.0826	583.8144	297.9481	194.3765	141.0155	1550.4463	2028.1352	2028.1352	197.9775	525.9808	746.7860	223.5734	956.1592
34.60	22.31	27.26	19.48	15.73	13.40	44.43	50.82	50.82	15.88	25.88	30.84	16.87	34.89
4.96	4.23	3.32	1.11	0.73	7.00	2.77	5.95	5.95	2.43	3.69	3.54	2.44	6.59
0.356935	0.229261	0.281777	0.191558	0.061156	0.112908	0.122796	0.128034	0.128034	0.196648	0.309782	0.187763	0.236561	0.623150
145.36635	114.89813	140.46412	72.36353	84.56149	56.52938	183.95666	357.14489	357.14489	119.08568	113.76556	148.68682	76.10549	142.33744

MWS 20	2184.52	1.000	1.000	0.000162	6177.99	3089.00	0.00000007	1543.7192	44.33	4.02	0.060861	468.42115
MWS 21	2799.37	1.000	1.000	0.002142	466.92	233.46	0.00000077	529.0756	25.95	3.44	0.366230	92.27297
MWS 22	3359.83	1.000	1.000	0.001492	670.29	335.15	0.00000044	1010.5787	35.87	0.80	0.056692	133.60606
MWS 23	2163.39	1.000	1.000	0.001696	589.46	294.73	0.00000078	167.9584	14.62	1.74	0.257865	65.49539
MWS 24	3590.06	1.000	1.000	0.001556	642.52	321.26	0.00000043	370.1083	21.71	3.74	0.459127	99.38378
MWS 25	5698.55	2.000	7.154	0.001672	598.26	299.13	0.0000008	612.4560	27.92	0.51	0.063517	96.53182
MWS 26	6627.31	0.500	0.699	0.001307	765.20	382.60	0.000000059	1907.5373	49.28	3.76	0.253530	235.47223
MWS 27	1326.64	1.000	1.000	0.001238	807.54	403.77	0.0000009	146.9065	13.68	0.43	0.045818	62.65726
MWS 28	21282.49	1.111	3.100	0.001346	743.03	371.51	0.00000089	634.7094	28.43	2.50	0.531609	144.97638
MWS 29	5275.33	0.500	1.176	0.001218	821.02	410.51	0.00000069	1521.5651	44.01	0.81	0.052374	188.63422
MWS 30	1574.33	1.000	1.000	0.001063	940.59	470.29	0.00000068	3455.7629	66.33	1.64	0.036148	368.25172
MWS 31	5973.56	1.000	1.155	0.001200	833.22	416.61	0.00000121	3011.9747	61.93	0.91	0.044406	324.66124
MWS 32	2376.35	2.000	5.437	0.001472	679.45	339.73	0.00000186	563.6647	26.79	0.65	0.051512	93.34960
MWS 33	1548.07	1.000	1.000	0.001098	910.82	455.41	0.00000071	594.3165	27.51	0.62	0.032937	105.32678

MWS 34	1529.55	1.000	1.000	0.001348	741.63	370.81	0.00000088	243.8166	17.62	0.43	0.039103	66.53864
MWS 35	1326.64	1.000	1.000	0.001238	807.54	403.77	0.000000093	146.9065	13.68	0.43	0.045818	62.65726
MWS 36	1584.93	1.000	1.000	0.001388	720.60	360.30	0.00000088	121.4687	12.44	0.19	0.024979	55.58045
MWS 37	42747.47	1.841	1.673	0.001421	703.52	351.76	0.00000130	646.7129	28.70	1.84	0.562887	149.52040

#### 4.1. Linear Parameters:

- 4.1.1. **Average length of stream (Lum):** It exhibits that total stream length decreases with increasing order. The total average stream length of all order segments in the Sobha watershed is 151926.09m and the highest average length is MWS 28 i.e., 21282.49 m and the lowest in MWS 18 i.e., 1235.10m.
- 4.1.2. **Bifurcation ratio** ( $\mathbf{R}_b$ ): Bifurcation shows a small range of variation for different regions or for different environment except where the powerful geological control dominates (Strahler, 1952 & 1957). The Bifurcation ratio at sub-watershed level has been carried out in GIS environment. In Sobha watershed it ranges from 0.500 to 2.000 and the mean  $\mathbf{R}_b$  in sub-watersheds is 1.053. The MWS 10, 15, 18, 24 and 30 are having highest  $\mathbf{R}_b$  which exhibits low permeability and structural control over it.

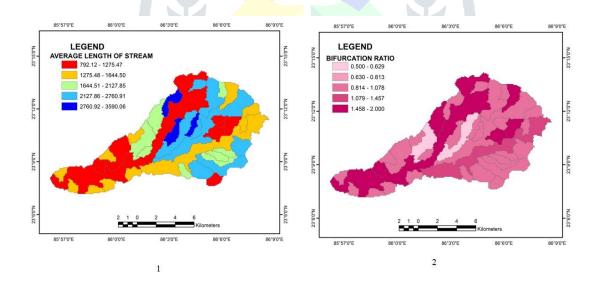


Figure 4: Linear Parameters-1. Average Length of Stream 2. Bifurcation Ratio

- 4.1.3. **Stream length ratio** (**L**<sub>ur</sub>): The stream length ratio of Sobha watershed ranges in 0.691 to 7.154 and changes of stream length ratio from one order to another indicating their late youth stage of geomorphic development. (Singh and Singh, 1997). The highest stream length ratio observed in MWS 24 and lowest in MWS 16 and 25.
- 4.1.4. **Stream frequency (F<sub>s</sub>):** generally high stream frequency is related to impermeable subsurface material, sparse vegetation, high relief, and low infiltration capacity of the region. The watershed has low stream frequency and it varies from sub-watershed to sub-watershed. The highest and lowest stream frequency occurred in MWS 30 and MWS 19 respectively.

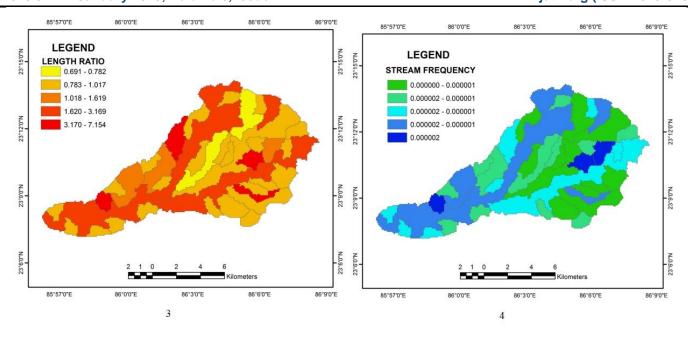


Figure 5: Linear Parameters-3. Length Ratio 4. Stream Frequency

#### 4.2. Aerial Parameters:

4.2.1. **Length of overland flow** (L<sub>g</sub>): It means the length of the runoff of the rain water on the ground surface before it is localized into definite channels. It is highest in MWS 19 and lowest in MWS 7.

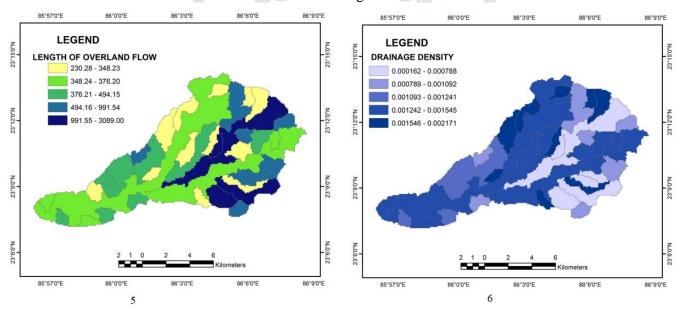


Figure 6: Aerial Parameters-5. Length of overland flow 6. Drainage Density

- 4.2.2. **Drainage density** ( $D_d$ ): It is vital element of drainage morphometry to study the landscape dissection, runoff potential, infiltration capacity of the land, climatic condition, and vegetation cover of the watershed. The drainage density of whole Sobha watershed indicates very low density which is only 0.04 m/m<sup>2</sup>. Sub-watershed 7, 20, 15 are having high drainage density resultant of weak or impermeable subsurface material, sparse vegetation, and rugged relief. Despite of this low drainage density in MWS 19, 11, 5 leads to coarse drainage texture and high permeable sub-soil material.
- 4.2.3. **Constant of channel maintenance (C):** It indicates the relative size of landform units in a drainage basin and has a specific genetic condition (Strahler, 1957). The total value of whole watershed is 34444.21 where MWS 7 has lowest value and MWS 19 has shown highest value.

4.2.4. **Form ratio** (**F**<sub>f</sub>): This is an important dimensionless property which enumerates the shape of the basin. Form factor of Sobha watershed is 40917.32 and it ranges from 121.469 to 4765.13 in subwatersheds. MWS 3 with highest form ratio showing high peak flow of shorter duration. Lowest form ratio observed in MWS 34. In short, the shape of the watershed is quite elongated.

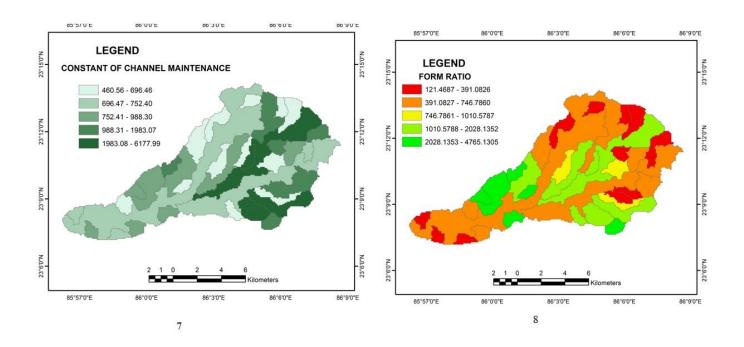


Figure 7: Aerial Parameters-7. Constant of channel maintenance 8. Form Ratio

4.2.5. **Elongation ratio** (R<sub>e</sub>): The elongation ratio runs between 12.44 and 77.89 for all subwatersheds. The lowest and highest elongation ratio reveals in MWS 36 and MWS 2 respectively. Sobha watershed is having high relief and steep valley side with almost circular in nature.

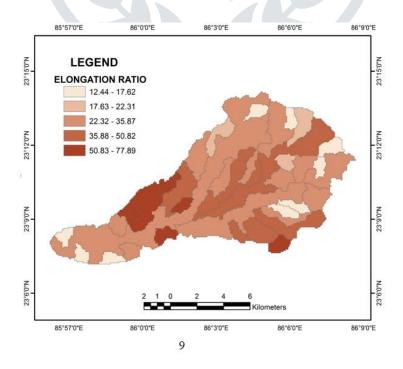


Figure 8: Aerial Parameters-9. Elongation Ratio

#### 4.3. **Relief Parameters:**

4.3.1. Relief ratio (R<sub>h</sub>): The elevation differences between the highest and lowest points on the valley floor of a sub-watershed. The relief ratio normally increases with decreasing drainage area and size of subwatersheds of a given drainage basin. It is highest in MWS 5 and lowest in MWS 36.

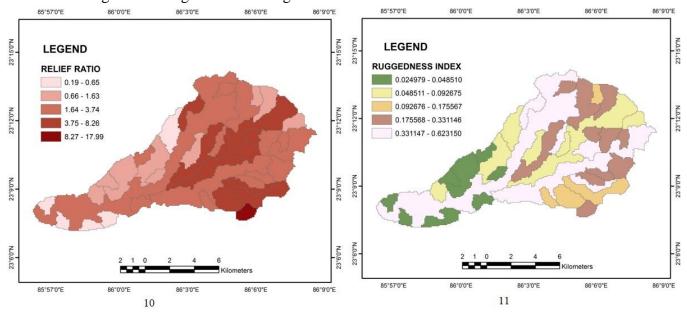


Figure 9: Relief Parameters-10. Relief Ratio 11. Ruggedness Index

4.3.2. Ruggedness number (R<sub>n</sub>): It is the product of the basin relief and the drainage density and usefully combines slope steepness with its length. The MWS 6, 3 7, 28 have highest value implies more prone soil erosion whereas MWS 36, 5, 33 have shown lowest value and indicates lower volume of soil erosion.

**Table 4:** Prioritization Classes based on Composite Score

Priority classes	Prioritized Zones (Composite score value)	MWS code numbers
Critical	55.58045 – 80.51194	7, 14, 16, 23, 34, 35, 36
Very high	80.51195 – 129.45791	4, 9, 10, 15, 18, 21, 24, 25, 32, 33
High	129.45792 – 225.54957	6, 8, 13, 17, 19, 22, 28, 29, 37
Moderate	225.54958 – 414.19855	1, 3, 11, 12, 26, 30, 31
Low	414.19856 – 570.29021	2, 5, 20

#### 5. Prioritization of Sub-Watersheds:

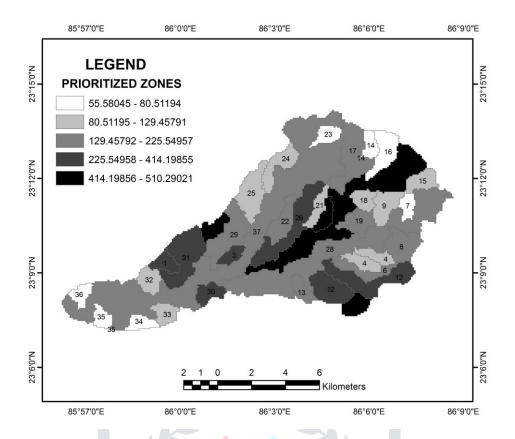


Figure 10: Prioritized zones of sub-watersheds

The composite parameter value of all 37 micro-watersheds of Sobha watershed is calculated and prioritization rating is shown in table- 4. The sub-watersheds have been classified into five priority zones according to their composite values.

Sub-watersheds falling under high to critical class are under very severe erosion susceptibility zone. Thus, need immediate attention to take up mechanical soil conservation measures gully control structures like check dams and grass waterways to protect the top soil loss. While sub-watersheds falling under moderate to low priority have very slight erosion susceptibility zone and may need agronomical measures to check the sheet and rill erosion.

#### **REFERENCES:**

- Ali, A. 2001. Macroeconomic variables as common pervasive risk factors and the empirical content of the Arbitrage Pricing Theory. Journal of Empirical finance, 5(3): 221–240.
- Binay, K. and Uday, K. 2012. Micro watershed characterization and prioritization using Geomatics technology for natural resources management. International Journal of Geomatics and Geosciences, 1 (4): 789-802.
- Biswas, S. Sudhakar, S. and Desai, V. R. 1999. Prioritization of sub watersheds based on morphometric analysis of drainage basin: A Remote Sensing and GIS Approach. Journal of Indian Society of Remote Sensing, 27 (3): 155-166.
- Black, P. E. 1972. Hydrograph responses to geomorphic model watershed characteristics and precipitation variables. Journal of Hydrology, 17(4): 309–329.

- Chopra, R. Dhiman, R. D. and Sharma, P. K. 2005. Morphometric analysis of sub-watersheds in Gurdaspur district, Punjab using remote sensing and GIS techniques. Journal of Indian Society of Remote Sensing, 33(4) 531–539.
- Chorley, R. and Morgan, M. 1962. Comparison of morphometric features, Unaka Mountains, Tennessee and North Carolina, and Dart moor, England. Geological Society of America Bulletin, 73(1): 17–34.
- Clarke, J. I. 1966. Morphometry from maps. Essays in Geomorphology, G. H. Dury, Ed., Elsevier, New York, NY, USA, 235–274.
- Das, S. Sudhakar, D. and Pardeshi, 2018. Integration of different influencing factors in GIS to delineate groundwater potential areas using IF and FR techniques: a study of Pravara basin, Maharashtra, India.
- DeCelles, P.G. Gehrels, G. E. Quade, J. and Ojha, T. P. 1998. Eocene early Miocene foreland basin development and the history of Himalayan thrusting, western and central Nepal. Tectonics, 17(5):741–765.
- DeCelles, P. G. Robinson, D. M. and Zandt, G. 2002. Implications of shortening in the Himalayan fold-thrust belt for uplift of the Tibetan Plateau. Tectonics, 21(6): 1–25.
- Diakakis, M. 2011. A method for flood hazard mapping based on basin morphometry: application in two catchments in Greece. Natural Hazards, 56 (3): 803–814.
- Faniran, A. 1968. The index of drainage intensity -a provisional new drainage factor. Australian Journal of Science, 31: 328–330.
- Horton, R. E. 1932. Drainage basin characteristics. Transactions of American Geophysics Union, 13: 350–361.
- Horton, R. E. 1945. Erosional development of streams and their drainage basins: hydro physical approach to quantitative morphology. Geological Society of America Bulletin, 56: 275–370.
- Jasmin, I. and Mallikarjuna, P. 2013. Morphometric analysis of Araniar river basin using remote sensing and geographical information system in the assessment of groundwater potential. Arab Journal of Geosciences, 6(10): 3683–3692.
- Javed, A. Khanday, M.Y. and Rais, S. 2011. Watershed prioritization using morphometric and land use/land cover parameters: a remote sensing and GIS based approach. Jour. Geol. Soc. India, 78: 63-75.
- Kumar, R. Ghosh, S. K. Mazari, R. K. and Sangode, S. J. 2003. Tectonic impact on the fluvial deposits of Plio-Pleistocene Himalayan foreland basin, India. Sedimentary Geology, 158(3-4): 209–234.
- McQuarrie, N. Robinson, D. Long, S. 2008. Preliminary stratigraphic and structural architecture of Bhutan: implications for the along strike architecture of the Himalayan system. Earth and Planetary Science Letters, 272(1-2): 105–117.
- Melton, M. A. 1957. An Analysis of the Relations among Elements of Climate, Surface Properties and Geomorphology. Columbia University, New York, NY, USA.
- Mishra, D. P. Dubey, and Tiwari, R. N. 2011. Morphometric analysis of Tons basin, Rewa District, Madhya Pradesh, based on watershed approach. Earth Science India, 4(3): 171–180.
- Mueller, J. E. 1968. An introduction to the hydraulic and topographic sinuosity indexes. Annals of the Association of American Geographers, 58(2): 371–385.

- Nag, S. K. 1998. Morphometric analysis using remote sensing techniques in the Chaka sub basins, Purulia district, west Bengal. Journal of Indian Society of Remote Sensing, 26: 69-76.
- Nag, S. K. and Chakraborty, S. 2003. Influence of rock types and structures in the development of drainage network in hard rock area. Journal of Indian Society of Remote Sensing, 31(1): 25–35.
- Panhalkar, S. S. and Pawar, C. T. 2011. Watershed Development Prioritization by applying WERM model and GIS techniques in Vedganga Basin (INDIA). ARPN Journal of Agriculture and Biological Science, 2: 38-44.
- Panhalkar, S.S. Mali, S. P. and Pawar, C. T. 2012. Morphometric Analysis and Watershed Development Prioritization of Hiranykeshi Basin in Maharashtra, India. International Journal of Environmental Science, 3: 525-534.
- Pareta, K. and Pareta, U. 2011. Quantitative Morphometric Analysis of a Yamuna Basin, India using ASTER (DEM) Data and GIS. International Journal of Geomatics and Geosciences, 2 (1): 248-269.
- Pareta, K. and Pareta, U. 2012. Quantitative geomorphological analysis of a watershed of Ravi River Basin, H.P. India. International Journal of Remote Sensing and GIS, 1(1): 41–56.
- Patton, P. C. and Baker, V. R. 1976. Morphometry and floods in small drainage basins subject to diverse hydrogeomorphic controls. Water Resources Research, 12(5): 941–952.
- Pearson, O. N. and DeCelles, P. G. 2005. Structural geology and regional tectonic significance of the Ramgarh thrust, Himalayan fold-thrust belt of Nepal. Tectonics, 24(4,): 1–26.
- Reddy, O. G. P. Maji, A. K. Gajbhiye, S. K. 2004. Drainage morphometry and its influence on landform characteristics in a basaltic terrain, Central India—a remote sensing and GIS approach. Int J Appl Earth Obs Geoinformatics, 6: 1–16.
- Romshoo, S. A. Bhat, S. A. and Rashid, I. 2012. Geoinformatics for assessing the morphometric control
  on hydrological response at watershed scale in the Upper Indus basin. Journal of Earth System Science,
  121(3): 659–686.
- Sachan, H. K. Kohn, M. J. Saxena, A. and Corrie, S. L. 2010. The Malari leucogranite, Garhwal Himalaya, Northern India: chemistry, age, and tectonic implications. Bulletin of the Geological Society of America, 122(11-12): 1865–1876.
- Schumm, S. A. 1956. Evolution of drainage systems and slopes in badlands at Perth Amboy, New Jersey. Bulletin of the Geological Society of America, 67: 597–646.
- Singh, S. Singh, M. B. 1997. Morphometric analysis of Kanhar river basin. Natl Geogr J India, 43(1): 31–43.
- Singh, P. Thakur, J. K. and Singh, U. C. 2012. Morphometric analysis of Morar River Basin, Madhya Pradesh, India, using remote sensing and GIS techniques. Environmental Earth Science, 68, (7): 1967–1977.
- Smart, J. S. and Surkan, A. J. 1967. The relation between mainstream length and area in drainage basins. Water Resources Research, 3 (4): 963–974.
- Srivastava, P. and Mitra, G. 1994. Thrust geometries and deep structure of the outer and lesser Himalaya, Kumaon and Garhwal (India): implications for evolution of the Himalayan fold-and thrust belt. Tectonics, 13(1): 89–109.

- Strahler, A. N. 1950. Equilibrium theory of erosional slopes approached by frequency distribution analysis. American Journal of Science, 248: 673-696, 800-814.
- Strahler, A. N. 1952. Hypsometric analysis of erosional topography. Bulletin of the Geological Society of America, 63: 1117–1142.
- Strahler, A.N. 1957. Quantitative analysis of watershed geomorphology. Transactions of American Geophysics Union, 38: 913–920.
- Strahler, A.N. 1964. Quantitative geomorphology of drainage basin and channel networks. Handbook of Applied Hydrology, V. T. Chow, Ed., McGraw Hill, New York, USA.
- Valdiya, K. S. 1980. Geology of the Kumaon Lesser Himalaya, Wadia Institute of Himalaya, Uttarakhand, India.
- Valdiya, K. S. 2001. Reactivation of terrane-defining boundary thrusts in central sector of the Himalaya: implications. Current Science, 81(11): 1418–1431.
- Vannay, J. C. and Grasemann, B. 2001. Himalayan inverted metamorphism and syn-convergence extension as a consequence of a general shear extrusion. Geological Magazine, 138(3): 253–276.
- Wakode, H. B. Dutta, D. Desai, V. R. Baier, K. and Azzam, R. 2011. Morphometric analysis of the upper catchment of Kosi River using GIS techniques. Arabian Journal of Geosciences, 6(2): 395 408.
- Williams, P. W. 1972. Morphometric analysis of polygonal karst in New Guinea. Geological Society of America Bulletin, 83 (3): 761–796.
- Yin, A. 2006. Cenozoic tectonic evolution of the Himalayan orogen as constrained by along-strike variation of structural geometry, exhumation history, and foreland sedimentation. Earth-Science Reviews, 76(1-2): 1–131.